PALAEOCENE—EARLY EOCENE BIOSTRATIGRAPHY IN NAREDA, SOUTHWESTERN KUTCH, WESTERN INDIA

K. K. TANDON†, V. K. MATHUR¹ AND R. K. SAXENA

GEOLOGY DEPARTMENT, LUCKNOW UNIVERSITY, LUCKNOW-226007

ABSTRACT

The present investigation has unravelled the presence of Palaeocene—Early Eocene sequence outcropping in the neighbourhood of village Nareda of southwestern Kutch. Micropalaeontological examination of the samples from the above sequence has revealed a rich assemblage of larger and smaller benthonic foraminifera and several easily identifiable planktonic foraminiferal taxa. Among planktonic taxa there are a few index species which seem to have significance in biostratigraphy. Aided by these index taxa, four planktonic foraminiferal zones have been recognised in the present succession and have been correlated on regional as well as inter-continental scale.

INTRODUCTION

The Palaeocene—Early Eocene succession under study is exposed in several nala sections near village Nareda (23°34'30": 68°41'30"), southwestern Kutch. The Deccan Trap in the area forms the barement, over which lie unconformably the Palaeocene rocks. These are in turn disconformably followed by Early Eocene rocks. The shales, clays and marls are the principal constituents of the sequence. Of the total 33.5 m thick succession, the Palaeocene is represented by 15 m thick sequence, whereas the strata representing the Early Eocene are 18.5 m in thickness. The succession as observed in the field is as follows:

Table 1

Yellow marl (3 m) Chocolate clay (2 m) Greenish grey, ferruginous marl (1 m) Early Erown clay (2.5 m) Eocene Glauconitic sandstone (2 m) Grey shale (2 m) Gypseous clay (3 m) Grey clay (3 m) ---Disconformity-Unfossiliferous clay (8 m) Venericardia shale (3 m) Carbonaceous shale (1 m) Palaeocene Pyritic clay (1 m) Grey clay (0.5 m) Carbonaceous shale (1.5 m)

Deccan Trap

The purpose of the present paper is to recognize the planktonic foraminiferal taxa and to arrive at a definite biostratigraphic unification, of course local in nature, of the above strata in the present area. The biozones defined here have been briefly described and correlated with other zones at regional level as well as those of the planktonic foraminiferal scheme at inter-continental level. From the historical review as presented below, it is probably true that for this part (Palaeocene—Early Eocene) of Tertiary succession in Kutch, none of the previous works, which were either for the most part the records of some fossil occurrences at respective localities of Kutch or were simply the descriptive accounts of fossils contained therein, ever dealt with biostratigraphic aspect of micropalaeontology.

PREVIOUS WORK

Biswas (1965) recognised Palaeocene of the mainland Kutch as the Madh Series and suggested its age on the basis of pollen and spore assemblage. Later, Biswas and Raju (1973) and Raju (1970) held the view that the Madh Series was constituted mainly by continental deposits. This contention was refuted by a report of some planktonic foraminifera from the Palaeocene rocks by Mathur et al. (1970). The occurrence of these planktonic species in the Palaeocene rocks established that these deposits were in part of marine nature and ranged in age from Early to Middle Palaeocene. The occurrence of marine Palaeocene rocks in Kutch was also recorded by Tandon (1971). Mathur (1972) described some nannofossils from Patcham Island and dated them as Palaeocene -Early Eocene. Samanta (1974), while making some biostratigraphic observations on the Palaeocene rocks of the India-Pakistan region, reviewed the earlier works

[†] Deceased Jan. 15, 1978

Present Address: Orissa Circle (South), G.S.I., Bhubaneshwar, Orissa.

on the rocks of this epoch in Kutch.

Tandon (1962) was first to record the presence of Early Eocene as fossiliferous Laki beds from Nareda, Kutch. Biswas (1965) referred to them as Kakdi Stage. Bhatt (1968) reported and described some planktonic foraminifera from these rocks.

FAUNAL ANALYSIS

The samples from the horizons of above succession were processed by boiling with sodium carbonate for several hours and washing over 50, 70, 100, 120, and 170 mesh sieve. The examination of the dried residue of the samples has brought to light the occurrence of microfossils in some of them. The faunal assemblage so obtained consists of ostracoda, larger and smaller benthonic foraminifera and planktonic foraminifera. The foraminiferal assemblage in the lower most 2 m thick sequence of carbonaceous shales and grey clays is represented by Globorotalia trinidadensis, Globorotalia perclara, Globorotalia compressa and Globorotalia spp. and also by several benthonic foraminiferal taxa. The following 2 m thick horizons represented by pyritic clays and carbonaceous shales are barren of microfauna. However, a change in the fauna is observed in the overlying 3 m thick band of shales of varying colours with the appearance of Globorotalia inconstans and disappearance of G. trinidadensis and G. compressa. Practically no microfauna could be recovered from the samples of further 8 m thick sequence of clays lying above the fossiliferous shales. They are rather characterized by pyritic inclusions which testify to the anaerobic conditions during deposition of this unfossiliferous band of clays.

The assemblage in the Early Eocene sequence too is rich in fauna and reveals fairly good number of both benthonic and planktonic foraminiferal taxa. In the lower 3 m thick sequence of sediments, the assemblage is dominated by the foraminiferal taxa like Globorotalia rex, G. wilcoxensis, Praeindicola bikanerensis and several species of Schackoinella. The accompanying larger foraminifera worked out by Tandon (1966) include Nummulites atacicus, N. mamilla, Assilina daviesi, A. spinosa, and A. subspinosa. The overlying 15.5 m thick sequence of beds is characterized by the absence of G. rex and presence of Globorotalia prolata, G. esnaensis and G. collactea. Among the larger foraminifera, all species of the underlying portion as mentioned above are present.

FORAMINIFERAL BIOZONATION

The following biozonation of the sequences (Palaeocene—Early Eocene) exposed at Nareda, southwestern Kutch has been proposed mainly based on the distribution of planktonic foraminifera. Four biozones of local nature are recognized, which are described below in ascending order (Table 2).

PALAEOCENE

Two biozones observed in this sequence are given below.

Globorotalia trinidadensis Zone: Thickness—2 m; Lithology—Carbonaceous shales and grey clays.

Characterisation: This zone contains an assemblage marked by Globorotalia trinidadensis, Globorotalia perclara, Globorotalia compressa and Globorotalia spp. The top of the zone is marked by the disappearance of G. trinidadensis and G. compressa. But G. perclara continues above.

Age: The zonal taxa and G. compressa which are known to occur in the basal part of Palaeocene suggest Early Palaeocene age to this zone.

Overlying the *G. trinidadensis* zone are unfossiliferous horizons of pyritic clays and carbonaceous shales, which are again overlain by fossiliferous grey coloured clays.

Globorotalia inconstans Zone: Thickness — 3 m; Lithology—Shales containing Venericardia sp. A.

Characterization: This zone is marked by local range of the zonal taxa which is accompanied by Globorotalia perclara of the preceding zone. Both become extinct at the top of this zone.

Age: With the help of the zonal taxa, this zone can be referred to the middle part of Palaeocene.

G. inconstans zone is followed by faunistically barren horizons doubtfully assignable to the later part of Palaeocene on account of their stratigraphic position. Overlying the unfossiliferous horizons is the Early Eocene sequence.

EARLY EOCENE

It is again divisible into two biozones.

Globorotalia rex Zone: Thickness—3 m; Lithology—Grey coloured plastic clays with a thick clayey limestone band near the base, at places small quantity of gypsum present.

Characterization: This zone is characterized by partial range of the zonal taxa; other planktonic foraminifera occurring in this zone are Globorotalia wilcoxensis, Globorotalia spp., Praeindicola bikanerensis, Globigerina sp. and Schackoinella spp., G. rex and G. wilcoxensis disappear at the top of this zone.

Age: The presence of zonal taxa is indicative of an early part of the Early Eocene.

Globorotalia prolata Zone: Thickness—15.5 m; Lithology—Gypseous brown clay, grey shale, glauconitic sandstone, brown clay, greenish grey marl, chocolate clay, yellow and brown coloured marl with a thin band of limestone at the top.

Characterization: Globorotalia prolata, Globorotalia esnaensis and Globorotalia collactea characterize this zone. Praeindicola bikanerensis, Schackoinella spp. and Globigerina sp. of the preceding zone are also present in this zone.

Table 2—Showing Lithounits and Biostratigraphic units of Palaeocene—Early Eocene sequence at Nareda, Southwestern Kutch.

Time U	lloit				
Epoch		Litholog	Lithounit	Bioz o n e	
w Z	_ge		Yellow Marl (3m)		
			Chocolate Clay (2m)		
С)	Ypresian		Greenish grey ferruginous Mart(Im)		
E 0 C			Brown Clay (2.5m)	Globorotalia prolata Zone	
			Glauconitic Sandstone (2m)		
		The column The	Grey Shale (2m)		
ARLY		XXXXXXXX XXXXXXX XXXXXXXX XXXXXXXX XXXXX			
Ш			Grey Clay (3 m)	Globorotalia rex Zone	
PALAEOCENE	Landenian		Unfossiliferous Clay (8m)	Faunistically barren Zome	
	Heersian		Venericardia Shale (3m)	Globorotalia inconstans Zone	
			Carbonaceous Shale (Im)	Faunistically barren Zone	
	Danían		Pyritic Clay (Im)	raumsticany dairem 2009	
			Grey Clay (0.5m) Carbonaceous Shale (1.5m)	Globorotalia trinidadensis Zone	

Age: G. prolata suggests that this zone is referable to the latter part of Early Eocene.

BIOSTRATIGRAPHIC CORRELATION

Authors have attempted the biostratigraphic correlation of the planktonic zones of the Palaeocene—Early Eocene succession of Nareda, southwestern Kutch with those of Rajasthan, Cambay basin, Cauvery basin, Andaman Islands, etc. in the Indian subcontinent and those of Trinidad, Italy, Egypt, U.S.S.R., etc. in the light of the correlation of Palaeocene foraminiferal zones suggested by Berggren (1971), El-Naggar (1969), Mohan and Pandey (1971) and Mohan et al. (1977).

Excepting Cauvery basin, Southern Shillong shelf and Andaman Islands, the biostratigraphic picture of the Palaeocene—Early Eocene is not very encouraging in most of the basins of the Indian subcontinent in view of poorly fossiliferous strata in the succession and lack of continuous sequence. The biostratigraphic correlation of the Palaeocene—Early Eocene strata in the Indian region has been attempted by many workers. The most convincing and pertinent to the present context is that of Raju (1970), Mohan and Pandey (1971) and Mohan et al. (1977).

Raju (1970) recognized the Palaeocene/Eocene boundary between Globorotalia pseudomenardii and Globorotalia velascoensis zones. On the basis of appearance of Nummulites/Assilina. According to him, in Cauvery basin where the presence of a seemingly continuous succession of Palaeocene—Eocene is found, Nummulites/Assilina appears at the base of Globorotalia velascoensis zone and can be taken to mark the onset of Palaeocene at the recommendation of Eocene Colloquium (Paris, 1968; Brabb, 1969). Subsequently, Mohan and Pandey (1971) introduced a little change in the boundary placement and suggested that the boundary be placed between Globorotalia pusilla pusilla zone and Globorotalia pseudomenardii

zone on account of the appearance of Nummulites/Assilina in lower part of Globorotalia pseudomenardii zone in Southern Shillong shelf. However, Mohan et al. (1977) share the opinion of Berggren (1971), so far as the placement of Palaeocene/Eocene boundary is concerned. Accordingly, the Pseudohastigerina appearance datum which coincides with the top of G. velascoensis zone (G. velascoensis—G. subbotinae zone of Berggren) marks the Palaeocene/Eocene boundary in the Indian basins, especially the Cauvery basin.

The lack of seemingly continuous succession and relatively poorly fossiliferous nature of horizons have actually marred the feasibility of complete biostratigraphy of the Palaeocene—Early Eocene succession in Kutch. However, the sequence at Nareda exposes some fossiliferous horizons, the microfauna of which renders the recognition of a few biostratigraphic zones practically possible. These biostratigraphic zones correlate well with those demarcated elsewhere for the Palaeocene—Early Eocene succession.

PALAEOCENE

Globorotalia trinidadensis zone of the present succession can be correlated with Globorotalia compressa/Globigerina daubjergensis zone of El-Naggar (1969), Globorotalia trinidadensis zone of Bolli (1957), Globorotalia trivalis subzone of Subbotina (1953, 1960), Globorotalia trinidadensis zone of Luterbacher (1964) and latter part of Globorotalia daubjergensis-Globigerina pseudobulloides zone of Berggren (1971) Fig. 2. This zone is referable to the latest part of the Danian Stage (Early Palaeocene). The Indian equivalents have been shown in Fig. 1. G. trinidadensis zone is followed by a faunistically barren zone which is again overlain by Globorotalia inconstans zone.

Globorotalia inconstans zone of the present succession correlates well with Globorotalia pusilla subzone (of Globo rotalia angulata zone) of El-Naggar (1969), Globorotalia

AGE	A	RAJAST HAN (Sigal & Singh;Unbub.in Mohan (Pandey,1971)	CAMBAY BASIN (Dutta et al, 1969)	CAUVERY BASIN (Raju, 1970)	ASSAM HIMALAYA (Mohan a Pandey,1971)	ANDAMAN ISLAND (Mohana Pandey, 1971)	KUTCH (Raju,1970)	KUTCH (Present Work)
EN EN		N atacicus		G palmerae Z		Data not Conclusive	Data not Conclusive	
EOC	:	G aragonensis Z.	Data Not Conclusive	G aragonensis Z	N. atabigina Ass	Garagonensis Ass	Data N atacicus G pro	G prolata Z
R X		G.formosa - A granulosa		G formosa formosa - G subbotinae Z		G formosa tormosa Ass.		Α
EA		Ass		G subbotinae-G aequa Z		Data not	GrexAss granulesa	G rex 2
E PORTO	(Umula)	Givelascoensis Z.	Nummulites Ass. 2	G velascoensis Z.			Data net Ass Cenclusive	Faunistically
CEN	(1 and	G pseudomenardii Z		G pseudomenardii Z	Nummulites - Discocyclina - M miscella Ass	M miscella Ass		barren
0 0	11.00	G pusilla-pusilla Z		G angulata 2	F primaeva-M miscella Ass.	Data not		G inconstant Z
A PAGE	(Ha	Data Nos	G pseudobulloides	G uncinata Z	Data not Conclusive	Conclusive		Faunisticily
A Q	Cu.	Conclusive	Ass	G trinidadensis Z	G pseudobulloides	G-pseudobulloides-		G.trinidadensis Z
P. P. L.	000		Deta not Conclusive	Data not Conclusive	S triloculinoides- Ass	Striloculinoides - Ass.		

Abbreviations-Ass.-Assemblage, A.-Assilina, F.-Fasciolites, G.-Globorotalia, M.-Miscellanea, N.-Nummulites, S.-Subbotina and Z.-Zone.

Fig. 1—Showing Inter-regional correlation of Palaeocene—Early Eocene rocks in India.

AGE		Bolli (1957)	EI-	Naggar (1969)	Subbotin	a (1953,60)	Luterbracher (1964)	Berggren (1971)	Present Work	
LU Z		G.palmerae Zone	Not studied		Zone of conical Globorotalids		G. bullbrooki Zone	A. densa Zone		
YEOCE	no	G.aragonensis Zone					G. aragonensis Zone	G aragonensis Zone	G. prolata Zone	
	presio	G.formosa formosa Zone					G. formosa formosa/	G formosa Zone		
	>	G.rex Zone	G.wilcoxensis Zone		b & C	G. margino-	G. subbotinae Zone	G. subbotinae - P.	G. rex Zone	
		G. Fex. 7. one			comp- subzone	G. aequa Zone	wilcoxensis Zone			
	Palse	G.velascoensis Zone	G.velascoensis Zone	G.aequa - G	- dopo -	globo - G. crassata/ A intermedia subzone a G. cone of otaliid- like subzone	G. velascoensis	G.velascoensis-G.subbotinae	Zone Faunistically barren	
				esnaensis subzon			Zone	G.velascoensis Zone		
		G.pseudomenardiiZone		G pseudomenar- dii subzone			G.pseudomenardii Zone	G.pseudomenardii Zone		
Z U	Mid. Palaeocene (Heersia)	G.pusilla pusillaZone	ata	G.pusilla subzone	a (Zone of		G.pusilla pusilla Zone	G. pusilla – G. angulata Zone	G. inconstans Zone	
E @ C			G. angula Zone	G.uncinata subzone	rotaliid- like globoro-		G. uncinata Zone	G.uncinata-G.spiralis Zone	Faunistically barres	
٦	•	G. trinidadensis Zone	1	mpressa/GI. taliids) ergensis Zone		G.trivialis subzane	G. trinidadensis Zone	Gl. daubjergensis - Gl.	G.trinidadensis Zone	
	Lr. Palaeoc (Danian)	Rzehakina epigona					GI.pseudobulloides/ GI.daubjergensisZone			
	٦	Z onule					Gleugubina Zone	pseudobulloides Zone		

Abbreviations-A.-Acarinina, G.-Globorotatia, Gl.-Globigerina and P.-Pseudohastigerina.

Fig. 2. Showing Inter-continental biostratigraphic correlation of Palaeocene-Eocene Epochs.

pusilla pusilla zone of Bolli (1957), latter part of Globorotalia inconstans subzone of Subbotina (1953, 1960),
Globorotalia pusilla pusilla zone of Luterbacher (1964) and
Globoratalia pusilla—Globorotalia angulata zone of Berggren
(1971) Fig. 2. The biostratigraphic position of this zone
suggests that it is referable to the latter part of Heersian
(Montian) stage (latter part of the Middle Palaeocene).
Its Indian equivalents have been shown in Fig. 1. The horizons overlying this zone have been found to be faunistically
barren and are in turn overlain by horizons showing
the presence of Globorotalia rex, an Early Eocene index
species. The faunistically barren horizons may possibly
be assigned to the Late Palaeocene by way of it being
intermediate in stratigraphic position between Middle
Palaeocene zone and Early Eocene zone.

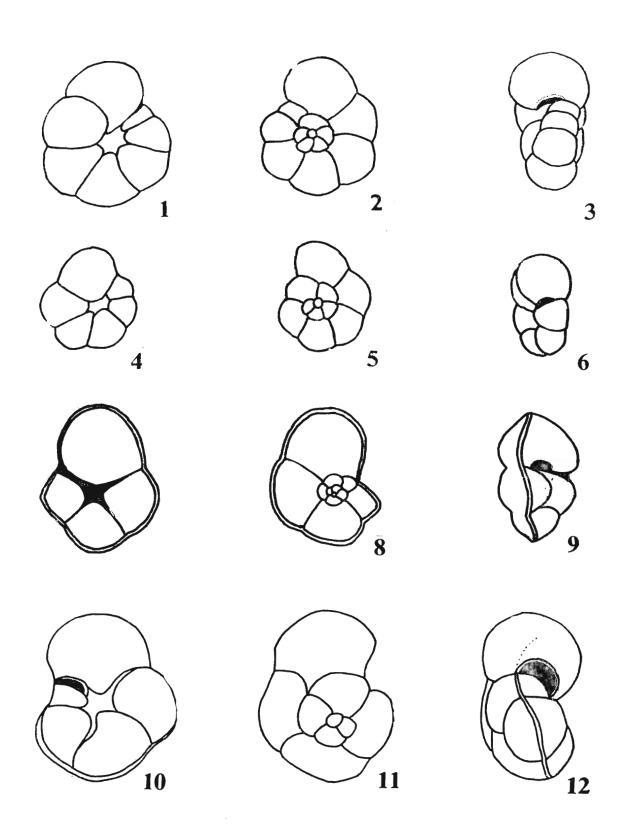
EARLY EOCENE

Globorotalia rex zone of the present study seems correlatable with Globorotalia wilcoxensis zone of El-Naggar (1969), Globorotalia marginodentata zone (C) of Subbotina (1953), combined Globorotalia aequa and early part of Globorotalia formosa formosa—Globorotalia subbotinae zones of Luterbacher (1964) and Globorotalia subbotinae—Pseudohastigerina wilcoxensis zone of Berggren (1971) Fig. 2. It refers to the early half of Ypresian stage (early part of Early Eocene). The equivalents of the Indian sub-

continent are shown in Fig. 1.

Globorotalia prolata zone at the top of the present succession under study appears to be in correspondence with the combined Globorotalia formosa formosa, Globorotalia aragonensis and Globorotalia palmerae zones of Bol (1957), lower part of the zone of conical Globorotalia (Sul hotina, 1960), combined latter part of Globorotalia formos formosa—Globorotalia subbotinae, Globorotalia aragonens and early part of Globorotalia bullbrooki zones of Lute bacher (1964) and the combined Globorotalia formos Globorotalia aragonensis and Acarinina densa zones of Berggre (1971) Fig. 2. Its other Indian equivalents are shown Fig 1. The overlying horizons are unfossiliferous and a assignable to the early part of the Middle Eocene becau of their conformable position underneath the fossilifero Middle Eocene strata.

The present study of the planktonic foraminife from Palaeocene—Early Eocene at Nareda, southweste Kutch suggests the possibility of the presence of for planktonic foraminiferal zones: Globorotolia trinidade zone, Globorotalia inconstans zone, Globorotalia rex zo and Globorotalia prolata zone. The biostratigraphic lir of these zones have led us to conclude that the Palacene—Early Eocene succession under study contains sediments belonging to the Upper Danian, Heer (Montian), Landenian and Ypresian Stages (Fig. 2).



ACKNOWLEDGEMENT

We sincerely thank Prof. S. N. Singh, Head of the Geology Department, Lucknow University for providing facilities. Financial help from U.G.C. and C.S.I.R., New Delhi is thankfully acknowledged to Junior authors respectively.

REFERENCES

- Beckmann, J. P., El-Heiny, J., Kerdany, M. T., Said, R., and Viotti, C. 1969. Standard planktonic zones in Egypt. In:

 Proc. First Internat. Conf. Plankt. Microf. (1967), Ed. Bronnimann, P. and Renz, H. H. 92-103.
- Bhatt, D. K. 1968. Planktonic foraminifera from the Lower Eocene sediments of Kutch, India. Bull. O.N.G.C. 4(2): 13-17.
- Biswas, S. K. 1965. A new classification of the Tertiary rocks of Kutch, Western India. Bull. Geol. Min. Met. Soc. India. 35: 1-6.
- Biswas, S. K. 1971. Note on the Geology of Kutch. Quart Jour. Geol. Min. Met. Soc. India. 43(4): 223-235.
- Biswas, S. K. and Raju, D. S. N. 1973. The rock—stratigraphic classification of the Tertiary sediments of Kutch. *Bull. O.N. G.C.* 10(1-2): 37-45.
- Berggren, W. A. 1965. Some problems of Palaeocene-Lower Eocene planktonic foraminiferal correlation. *Micropal.* 11 (3): 278-301.
- Berggren, W. A. 1971. Multiple phylogenetic zonation of the cenozoic based on planktonic foraminifrea. *Proc. II Plankt. Conf. Roma* (1970), Ed. Farnacci, A. 1: 41-65.
- Bolli, H. M. 1957a. The genera Globigerina and Globorotalia in the Palaeocene—Lower Eocene Lizards Springs Formation of Trinidad, B. W. I. U. S. Nat. Mus. Bull. 215: 61-68.
- BOLLI, H.M. 1957b. Planktonic foraminifera from the Eocene Navet and San Fernando Formations of Trinidad, B. W. I. U.S. Nat. Mus. Bull. 215: 155-172.
- BOLLI, H. M. AND CITA, M. B. 1960. Upper Cretaceous and Lower Tertiary planktonic foraminifera from Poderno-d' Adda section, northern Italy. Internat. Geol. Congr. 21st session, Copenhagen, 1960, Sect. 5: 150-161.
- Brabb, E. E. 1969. Eocene Colloquium: Geotimes. 14(3): 19-20. EL-Naggar, Z. R. 1969. Correlation of the various planktonic foraminiferal zonations of the Palaeocene. In: Proc. Ist. Conf.

- Plankt. Microf. (1967). Ed. Bronnimann, P. and Renz, H. H.: 92-103
- LUTERBACHER, H. P. 1964. Studies in some Globorotalia from the Palaeocene and Lower Eocene of the Central Appennines. Eclog. Geol. Helv. 57(2):631-730.
- Mathur, Y. K. 1972. Plant fossils from the Kuar Bet Patcham Island, Kutch. Curr. Sci. 41: 488-489.
- MATHUR, Y.K., SOODAN, K.S. MATHUR, KANWAL, BHATIA, M.L., JUYAL, N. P. AND PANT, J. 1970. Microfossil evidences on the presence of Upper Cretzceous and Palaeocene sediments in Kutch. Bull. O.N.G.C. 7 (2): 104-114.
- МОНАН, M. AND PANDEY, J. 1971. Foraminiferal control in the 'Геттіату basins of India. Jour. Pal. Soc. India. 16: 78-88.
- Mohan, M., Kumar, P. and Narayanan, V. 1977. Palaeocene— Eocene Succession in Cauvery basin. Jour. Geol. Soc. India. 18(8): 401-411.
- RAJU, D. S. N. 1970. Zonal distribution of selected foraminifera in the Cretaceous and Cenozoic sediments of Cauvery Basin and some problems of Indian biostratigraphic classification. *Publ. Gent. Adv. Stud. Geol. Panjab Univ. Chandigarh.* 8: 85-110.
- Samanta, B. K. 1973. Planktonic foraminifera from the Palaeocene-Eocene succession in the Ranikot Nala, Sulaiman Range, Pakisten. Bull. Brit. Mus. Nat. Hist., Geol. 22(6): 421-482.
- Samanta B. K. 1974. The limits and subdivision of the Palaeocene with remarks on the marine occurrences recorded in the India-Pakistan region. *Geol. Min. Met. Soc. India*, Golden Jubilee. 46: 183-205.
- SINGH, S. N. 1971. Planktonic foraminifera in the Eocene Stratigraphy of Rajasthan, India. Proc. II Conf. Plankt. Microf. (Roma, 1970). II: 1169-1181.
- Subbotina, N. N. 1953. Globigerinidae, Hantkeninidae and Globorotaliidae: Fossil foraminifera of the U.S.S.R. (VNIGRI), Trudy, New Ser. (76): 1-296.
- Subbotina, N. N. 1960. Palaeogene foraminifera of the Palaeogene deposits of Southern U.S.S.R.: In Menner (editor): "Palaeogene deposits of the Southern European part of the U.S.S.R."

 Academy of Science, U.S.S.R., Press, Moscow. 24-36.
- Tandon, K. K. 1962. Fossiliferous Laki beds from Kutch, India. Gurr. Sci. 31:65-66.
- Tandon, K. K. 1966. Stratigraphy and Micropaleontology of the Tertiary rocks of S. W. Kutch. Ph.D. Thesis (unpublished) Luckuow University, Lucknow.
- TANDON K. K. 1971. Occurrence of Venericardia beaumonti d'Archaic and Haime from Nareda, S. W. Kutch, India. Geophytology. 1 (1): 70-74.

EXPLANATION OF PLATE

PLATE I

- 1-3. Globorotalia trinidadensis 1. Umbilical view; 2. Spiral view; 3. Apertural view, ×150.
- 4—6. Globorotalia inconstans 4. Umbilical view; 5. Spiral view; 6. Apertural view, ×150.
- 7-9. Globorotalia rex 7. Umbslical view; 8. Spiral view; 9. Apertural view, ×150.
- 10-12. Globorotalia prolata 10. Umbilical view; 11. Spiral view; 12. Apertural view, ×150.